

A decision support tool to estimate unregulated, daily streamflow at ungauged sites in the Connecticut River Basin, northeast United States

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1 **Abstract.**

2 Streamflow information is needed for any number of hydrologic applications. Because most
3 stream reaches are ungauged, this information is commonly needed for rivers that have no
4 readily available measurements of streamflow. In the Connecticut River Basin, dam operation
5 and its effects on the aquatic habitat are of particular interest. Here, daily streamflow is needed
6 for use as input to dam simulation and optimization models as well as to develop ecological-flow
7 prescriptions for rivers and streams. To provide a common scientific foundation for water
8 allocation decisions, a freely available and easy-to-use software tool termed the Connecticut
9 River UnImpacted Streamflow Estimation (CRUISE) tool was developed to estimate a daily
10 streamflow time series at ungauged locations in the Connecticut River Basin. CRUISE is used in
11 conjunction with the U.S. Geological Survey StreamStats web application. Through the
12 coupling of CRUISE and StreamStats, users are able to point and click on a stream location of
13 interest and obtain a delineated catchment as well as a daily time series of streamflow. Daily
14 streamflow was shown to be reliably estimated by the CRUISE tool, with efficiency values
15 between the observed and estimated streamflows ranging from 0.69 to 0.92 and ecologically-
16 relevant streamflow statistics derived from the estimated daily streamflow to be generally within
17 +/- 10 percent of the streamflow statistics computed from the observed daily streamflow values.

18 **Keywords:** decision support; ungauged; unaged; streamflow; water availability; basin
19 delineation; water resources

20 1. Introduction

21 Streamflow information at ungauged stream reaches is needed for any number of hydrologic
22 applications. Furthermore, when streamflow information is presented in easy-to-use, freely-
23 available software tools, this information can provide a scientific framework for water-allocation
24 negotiation amongst stakeholders. Unfortunately, many rivers of the world are not gauged and
25 streamflow information is not available where it is often needed to make informed water-
26 allocation decisions. Additionally, there has been increasing emphasis on the need for daily
27 streamflow time series to understand the response of ecology to river regulation and develop
28 streamflow prescriptions to restore and protect aquatic habitat [*Poff et al.*, 1997]. For these
29 reasons, a software tool was developed to estimate daily streamflow time series at ungauged
30 streams in the Connecticut River Basin (CRB), located in the northeast United States. The
31 Connecticut River UnImpacted Streamflow Estimation (CRUISE) software tool is based on a
32 geographic information system (GIS) that allows users to point and click on an ungauged stream
33 location of interest in the CRB. The tool then delineates a contributing area to the stream location
34 and estimates a daily streamflow time series.

35 The CRB has thousands of dams along the mainstem and tributary rivers that are used for
36 hydropower, flood control, and water supply just as the CRB is home to a number of important
37 fish species that rely on the river for all or part of their life cycle. These competing interests for
38 water led the Army Corps of Engineers and The Nature Conservancy to embark on a partnership
39 to understand how dam management can be optimized to meet both human and ecological needs
40 for water. To answer this question, daily streamflow time series are needed at locations in the
41 CRB that have ecological constraints on water (locations where important or protected fish or
42 ecological communities reside or rely on for life), human constraints on water (locations on the

43 river that are dammed or otherwise managed), or have both constraints to consider. Often times,
44 these locations are unmonitored.

45 Methods to estimate daily streamflow time series at ungauged locations can be broadly
46 characterized under the topic of regionalization [*Blöschl and Sivapalan, 1995*], an approach
47 which pools information about streamgauges in a region and transfers this information to an
48 ungauged location. Generally there are two main categories of information that is pooled and
49 transferred: 1) rainfall-runoff model parameters that are calibrated at gauged catchments and
50 transferred in some way to an ungauged location [see *Zhang and Chiew, 2009* for a review] and
51 2) gauged streamflows, or related streamflow properties, are directly transferred to ungauged
52 locations. Examples of this type of regionalization approach include geostatistical methods such
53 as top-kriging [*Skøien and Blöschl, 2007*] and more commonly used methods such as the
54 drainage-area ratio method as described in *Archfield and Vogel [2010]*, the MOVE methods
55 [*Hirsch, 1979*], which are primarily used to patch and extend missing daily streamflow
56 information in existing records but can be easily extended to ungauged methods, and a non-linear
57 spatial interpolation method, applied by *Fennessey [1994]*, *Hughes and Smakhtin [1996]*,
58 *Smakhtin [1999]*, *Mohamoud [2008]*, and *Archfield et al. [2010]*. For the software tool presented
59 in this paper, a hybrid approach combining the drainage-area ratio and non-linear spatial
60 interpolation methods is used to estimate daily streamflow time series.

61 Software tools to provide streamflow time series at ungauged locations have been previously
62 published for fixed catchments. *Smakhtin and Eriyagama [2008]* and *Holtzschlag [2009]*
63 introduced software tools to provide monthly streamflows for ecological streamflow assessments
64 around the globe and in the Great Lakes region of the United States, respectively. *Williamson et*
65 *al. [2009]* developed The Water Availability Tool for Environmental Resources (WATER) to

66 serve daily streamflow information at fixed stream locations in non-karst areas of Kentucky.
67 These existing tools provide valuable streamflow information, and in most cases at the monthly –
68 not daily – time step for fixed catchments; yet, often the locations of most interest on a river are
69 not coincident with the pre-defined hydrologic units presented in these tools. There are few – if
70 any – software tools that can provide daily streamflow for user-specified (unfixed) locations
71 within a region. The U.S. Geological Survey StreamStats tool [*Ries and others*, 2008] provides
72 the utility to delineate a contributing area to a user-selected location on a river; however, only
73 streamflow statistics – not streamflow time series – are provided for the ungauged location.

74 The CRUISE tool is one of the first such tools to combine the utility of catchment delineation
75 at any location along a stream with the estimation and serving of daily streamflow information.
76 *Archfield et al.* [2010] developed a GIS-based software tool to estimate daily streamflow;
77 however, this tool requires software and licensing not available to all users, and covers only the
78 state of Massachusetts. This paper extends the work of *Archfield et al.* [2010] and presents the
79 first such software tool to obtain daily streamflow time series at ungauged locations in a regional
80 framework that requires only the use of an internet connection and a fairly ubiquitous
81 spreadsheet program. This framework has the potential to be applied to other regions to provide
82 daily streamflow information for ungauged locations.

83 This paper first describes the study area and the data required by the software tool. The
84 underlying methods to estimate daily streamflow time series in the software tool are then
85 presented and the software tool and functionality are described. Lastly the utility of the software
86 tool to provide reliable estimates of daily streamflow is demonstrated.

87 **2. Study area and data**

88 The study area is located in the northeast United States and covers an area of
89 approximately 29,000 km² (fig. 1). The region is characterized by a temperate climate with
90 distinct seasons. Snowfall is common from December through March, with generally more
91 snowfall falling in the northern portion of the CRB. The geology and hydrology of the study
92 region is heavily affected by the growth and retreat of glaciers during the last ice age, which
93 formed the present-day stream network and drainage patterns [Armstrong *et al.*, 2008]. The
94 retreat of the glaciers filled the river valleys with outwash sands and gravel as well as fine- to
95 coarse-grained lake deposits [Armstrong *et al.*, 2008], and these sand and gravel deposits have
96 been found to be important controls on the magnitude and timing base flows in the southern
97 portion of the study region [Ries and Friesz, 2000].

98 Data from streamgauges located within the CRB and surrounding area are used in the
99 CRUISE tool to estimate daily streamflow time series at ungauged locations (fig. 1; table 1). The
100 63 study streamgauges have at least 20 years of daily streamflow record and have minimal
101 regulation in the contributing catchment to the streamgauge [Armstrong *et al.*, 2008; Falcone *et*
102 *al.*, 2010]. Previous work in the southern portion of the study area by Archfield *et al.* [2010]
103 showed that the contributing area to the streamgauge, percent of the contributing area with
104 surficial sand and gravel deposits, and mean annual precipitation values for the contributing area
105 are important variables in modeling streamflows at ungauged locations. For this reason, these
106 characteristics were summarized for the study streamgauges (fig. 2) and used in the streamflow
107 estimation process. Contributing area to the study streamgauges ranges from 0.5 km² to 1,845
108 km² with a median value of 200 km² (fig. 2A). Mean annual precipitation ranges from 101 cm
109 per year to 157 cm per year with a median value of 122 cm per year (fig. 2B). Percent of the

110 contributing area with surficial sand and gravel ranges from 0 percent to 67 percent with a
111 median value of 9.5 percent (fig. 2C).

112 **3. Methods underlying the software tool**

113 Streamflow in the CRUISE tool is estimated for a 44-year (16,071-day) period of record
114 spanning October 1, 1960 through September 30, 2004 using information from an index
115 streamgauge and catchment characteristics computed for the contributing area to the ungauged
116 stream location of interest (fig. 3). Catchment characteristics and the selected index streamgauge
117 are used to first estimate a continuous, daily flow-duration curve (FDC) for the 44-year
118 simulation period (fig 3). The estimated FDC at the ungauged location is then transformed to a
119 time series of streamflow values by the index streamgauge (fig. 3). The methods to estimate the
120 FDC, select the index streamgauge, and transform the FDC to a time series of daily streamflow
121 are explained in detail in the following sections.

122 **3.1 Estimation of the period-of-record flow-duration curve**

123 Estimation of the period-of-record FDC at an ungauged location remains an outstanding
124 challenge in hydrology. *Castellarin et al.* [2004] provides a review of several methods to
125 estimate FDCs at ungauged locations and found that no particular method was consistently
126 reliable. For this study, an empirical, piece-wise approach to estimate the period-of-record FDC
127 is used in the CRUISE tool (fig. 4). This overall approach is similar to that used by *Mohamoud*
128 [2008] and *Archfield et al* [2010] in that the FDC is estimated by first developing regional
129 regressions relating catchment characteristics to selected FDC quantiles and then interpolating
130 between those quantiles to obtain a continuous FDC.

131 Streamgauges having at least 20 years of daily streamflow record that also contain the
132 drought of record for the study area were used for this portion of the CRUISE tool development.
133 A total of 52 streamgauges fit these criteria (fig. 1). Streamflows at each of these streamgauges
134 were ranked and corresponding exceedence probabilities were determined using the Weibull
135 plotting position [*Stedinger et al.*, 1993]. Selected streamflow quantiles were then determined by
136 applying equation 2 presented in *Vogel and Fennessey* [1994].

137 With the exception of streamflows having less than or equal to a 0.01 probability of being
138 exceeded (streamflows with a probability of being exceeded more than 1 percent of the time),
139 selected quantiles on the FDC are estimated from explanatory variables (fig. 4) and a continuous
140 FDC is log-linearly interpolated between these quantiles to obtain a continuous FDC (fig. 4).
141 Relations between streamflow quantiles at the 0.02, 0.05, 0.1, 0.15, 0.2, 0.25, 0.3, 0.4, 0.5, 0.6,
142 0.7, 0.75, 0.8 and 0.85 exceedence probabilities were estimated by independently regressing each
143 streamflow quantile against catchment characteristics (table 2; fig. 4). Following the approach in
144 *Archfield et al.* [2010], relations between streamflow quantiles at the 0.9, 0.95, 0.98, 0.99 and
145 0.999938 were estimated by regressing streamflows at these quantiles against one another and
146 using these relations to recursively estimate streamflows (table 3; fig. 4). Recursively estimating
147 low streamflows, as was done in *Archfield et al.* [2010], exploits the strong structural relation
148 between the streamflow quantiles (fig. 4) and enforces the constraint that streamflows must
149 decrease as the exceedence probability increases. *Mohamoud* [2008] and *Archfield et al.* [2010]
150 observed that when regression is done against catchment characteristics, there is increased
151 potential for the estimated quantiles to violate the constraint that streamflows must decrease as
152 the exceedence probability increases because the uncertainty in the flow estimates is greatest at
153 the lowest portion of the FDC. Regressing quantiles against one another ensures that this

154 constraint is not violated. This is an alternative approach to that used by *Mohamoud* [2008], who
155 suggested discarding any estimated quantiles that violate the constraint.

156 All regressions were fit using the natural logarithms of the dependent and independent
157 variables. Retransformation was needed to obtain estimated streamflows in their original units of
158 measure. The Smearing adjustment [*Duan*, 1983] was used to eliminate the bias in the streamflow
159 estimates caused by the retransformation. Following the regression-screening protocol described
160 in *Archfield et al.* [2010], independent variables having coefficient that were statistically
161 significant at the 0.05 level were included in the final equations. For regression equations with
162 multiple significant independent variables, multicollinearity was evaluated using the variance
163 inflation factor [*Stedinger et al.*, 1993]. Independent variables with a variance inflation factor
164 higher than 2.5 were removed from the regression equations. Residuals were evaluated for
165 normality and streamgauges that exerted a high degree of influence on the regression were
166 omitted. Percent root-mean-square error and the Nash-Sutcliffe efficiency value [*Nash and*
167 *Sutcliffe*, 1970] were computed from a leave-one-out cross validation of each site used in the
168 development of the regression equations. Goodness-of-fit metrics, regression diagnostics and
169 coefficients, as well as explanatory variables for each quantile, are shown in tables 2 and 3.

170 *Archfield et al.* [2010] showed that estimated streamflows determined by log-linear
171 interpolation between exceedence probabilities of 0.01 or less do not match the shape of the FDC
172 in this range and this interpolation method creates a bias in the estimated streamflows, which can
173 substantially overestimate the peak streamflows. The shape of the FDC at the highest
174 streamflows is so complex that, instead of using another interpolation method, the CRUISE tool
175 uses scaled streamflows from an index streamgauge to estimate the highest streamflows at the
176 ungauged location. The assumption here is that the shape of the left tail of the FDC is better

177 approximated by the streamflow quantiles at an index streamgauge than by a curve fit. Therefore,
 178 for streamflows having less than or equal to a 0.01 probability of being exceeded, streamflows
 179 are scaled by a drainage-area approach (eqn. 1) in conjunction with the selected index
 180 streamgauge:

$$q_{p_u} = \frac{A_u}{A_g} q_{p_g} \quad (1)$$

181 where q_{p_u} is the value of the streamflow quantile at the ungauged location for
 182 exceedence probability, p , A_u is the contributing drainage area to the ungauged location, A_g is the
 183 contributing drainage area to the index streamgauge, and q_{p_g} is the value of the streamflow
 184 quantile at the index streamgauge for exceedence probability, p .

185 **3.2 Selection of the index streamgauge**

186 As shown in figure 3, the index streamgauge is used for two purposes in the CRUISE
 187 tool: 1) to estimate streamflows that have less than a 1-percent chance of being exceeded, and 2)
 188 to transform the estimated FDC into a time series of streamflow at the ungauged location. The
 189 index streamgauge is selected by the map-correlation method [Archfield and Vogel, 2010]. The
 190 map-correlation method selects the index streamgauge estimated to have the highest cross-
 191 correlation between streamflow time series at the index streamgauge and the ungauged location.
 192 Archfield and Vogel [2010] showed that the selection of the index streamgauge using cross-
 193 correlation between streamflow time series outperformed the selection of the nearest index
 194 streamgauge when used with the drainage-area ratio method to estimate daily streamflow time
 195 series at ungauged locations. This finding supports the use of the map-correlation method in the
 196 CRUISE tool for two reasons: 1) the drainage-area ratio approach is also used in the CRUISE

197 tool to estimate streamflows that have less than a 1-percent chance of being exceeded, and 2)
198 because the streamflow time series in the CRUISE tool is constructed by transferring the timing
199 of the streamflows at an index streamgauge to the ungauged location, it follows that one would
200 seek to select the index streamgauge that maximizes the cross-correlation between the
201 streamflows at the ungauged location and the index streamgauge.

202 Underlying the map-correlation method is a set of variogram models – one for each index
203 streamgauge – that is fitted to the observed cross-correlations between the streamflows at the
204 index streamgauge and each of the other index streamgauges in the study region. The map-
205 correlation method uses these variogram models with ordinary kriging to estimate the cross-
206 correlation between each index streamgauge and the ungauged location and, ultimately, selects
207 the index streamgauge whose streamflows are estimated to be most correlated with the ungauged
208 location of interest.

209 To develop the variogram models, observed cross-correlations between daily streamflows
210 were computed from a long-term, 30-year common period of record from January 1, 1960
211 through December 31, 1982 that was available at 45 of the study streamgauges (fig. 1). This 30-
212 year common period was selected to maximize the number of index streamgauges in the
213 CRUISE tool while ensuring the observed cross-correlations were not affected by small or
214 uneven sample sizes. The nonparametric, rank-based Kendall tau correlation measure was used
215 the estimate cross-correlation. Kendall tau estimates the monotonic relation between two
216 variables [*Helsel and Hirsch, 2002*] and, therefore, its application requires fewer assumptions
217 than the Pearson r correlation coefficient, which measures only the linear correlation between
218 two variables.

219 For each index streamgauge, the observed cross-correlation between that streamgauge
220 and each of the other streamgauges is determined. Then, the differences between cross-
221 correlation, formally expressed as the semi-variance [Isaaks and Srivastava, 1990], taken for
222 each pair of streamgauges were plotted against the separation distance between each of pair of
223 streamgauges. Therefore, for each of the 45 index streamgauges, this results in a plot of $\binom{44}{2}$
224 points. To discern a relation between the semi-variance and separation distance, the points were
225 placed into 12 bins, with each bin having a length of 200,000 meters. A spherical variogram was
226 fit to the binned values to provide a continuous relation between the semi-variance and
227 separation distance exactly as described by Archfield and Vogel [2010] and using the geoR
228 statistical package [Ribeiro Jr. and Diggle, 2001]. Variogram models were developed for each of
229 the 45 index streamgauges (table 4) and a leave-one-out cross validation procedure was applied
230 to evaluate the utility of the variogram model to estimate the cross-correlation between the
231 removed site and the index streamgauge upon which the variogram model was fit. The root-
232 mean-square error resulting from this cross validation at each index streamgauge is reported in
233 table 4.

234 3.3 Generation of streamflow time series

235 With an index streamgauge and estimated daily FDC at the ungauged location, a time
236 series of daily streamflow for the 44-year simulation period is then constructed by use of the
237 QPPQ transform method [Fennessey, 1994; Hughes and Smakhtin [1996]; Smakhtin, 1999;
238 Mohamoud, 2008; Archfield et al. 2010]. The term QPPQ-transform method was coined by
239 Fennessey [1994]; however, this method has been published by Smakhtin [1999], Mohamoud
240 [2008], and Archfield et al. [2010] under names including “non-linear spatial interpolation

241 technique” [Hughes and Smakhtin [1996]; Smakhtin, 1999] and “reshuffling procedure”
242 [Mohamoud, 2008]. The method assumes that the exceedence probability associated with a
243 streamflow on a given day at the index streamgauge also occurred on the same day as the
244 ungauged location. For example, if the streamflow on October, 1, 1974 was at the 0.9
245 exceedence probability at the index streamgauge, then it is assumed that the streamflow on that
246 day at the ungauged location also was at the 0.9 exceedence probability . To implement the
247 QPPQ-transform method in the CRUISE tool, a FDC is constructed for the observed streamflows
248 at the index streamgauge, keeping track of the dates associated with each exceedence probability.
249 The exceedence probabilities are then equated between the index streamgauge FDC and the
250 estimated FDC at the ungauged location. The date associated with each exceedence probability at
251 the index streamgauge is then transferred the estimated FDC at the ungauged location.

252 The QPPQ-transform method requires that each index streamgauge has streamflow
253 values for each day of the simulation period. Recall that the 45 index streamgauges shared a
254 common period of observed daily streamflow record for 30 years of the 44-year simulation
255 period. However, for 20 of the 45 index streamgauges, the observed streamflow record does not
256 cover the full 44-year simulation period (table 1). For these streamgauges, the MOVE3 [Vogel
257 and Stedinger, 1985] record extension method was utilized to estimate a complete 44-year period
258 of record. The software program Streamflow Record Extension Facilitator (SREF) [Granato,
259 2009] was used to extend the streamflow records. The QPPQ-transform method uses only the
260 timing of the streamflows at the index streamgauge and not the magnitudes of the streamflows;
261 therefore, the MOVE3-estimated streamflow values themselves are not used in the transfer
262 process.

263 **4. The Connecticut River UnImpacted Streamflow Estimator (CRUISE) software tool**

264 The Connecticut River UnImpacted Streamflow Estimator (CRUISE) tool is freely
265 available for download at <http://webdmamrl.er.usgs.gov/s1/sarch/ctrtool/index.html>. The
266 CRUISE tool website contains additional information about the software required, a user manual,
267 history of updates, file sizes, and contact information. To use the CRUISE tool, users must have
268 an internet connection, a web browser program, and Microsoft Excel version 2003 or higher. The
269 U.S. Geological Survey StreamStats tool [*Ries and others*, 2008] is first used to delineate the
270 contributing area to the ungauged location and compute the catchment characteristics needed to
271 estimate the FDC, and then the CRUISE tool, which is a customized Microsoft Excel spreadsheet
272 with Visual Basic macros is used to select the index streamgauge and compute the unimpacted
273 daily streamflow time series for the ungauged location.

274 The StreamStats tool operates within a web browser, and is accessible at
275 <http://streamstats.usgs.gov>. The StreamStats home page provides a general description of the
276 application. A gray box on the left side of the page contains a series of links to pages that
277 document how to use the application, define terminology, and so forth. Selecting the *Access*
278 *User Interface* link will cause a new browser window to appear in which the StreamStats user
279 interface will display a map of the United States. Selecting *Region* from the *Zoom To* pull-down
280 list above the map will cause a small window to appear. Selecting *Connecticut River SYE* from
281 the pull-down list in that window will cause the map in the StreamStats user interface to display
282 the extent of the CRB and also will cause an introductory page to appear that explains the
283 StreamStats functionality that is available for this area, provides citations to relevant reports, and
284 identifies other organizations that contributed to the application development.

285 The map navigation tools provided in the StreamStats user interface should be used to
286 locate a point along the stream of interest. With the map zoomed into a scale of at least 1:24,000,

287 pressing on the *Watershed Delineation* button, and then on the map at location of interest will
288 cause the catchment boundary for the selected location to be delineated and displayed on the map
289 (fig. 5A). Once the catchment is delineated, pressing on the *Basin Characteristics* button will
290 result in the appearance of a new browser window that contains a table of the catchment
291 characteristics for the selected location (fig. 5B). StreamStats uses the processes described by
292 ESRI, Inc. (2009) for catchment delineation and computation of catchment characteristics.

293 StreamStats provides a *Download* tool to export a shapefile of the contributing catchment
294 (fig. 5A) for use in other mapping applications. If catchment characteristics are determined
295 before the shapefile is created, then the catchment characteristics will be saved as attributes with
296 the shapefile.

297 The CRUISE tool consists of a Microsoft Excel spreadsheet with five worksheets. The
298 features of the CRUISE tool are shown in figure 5. The spreadsheet opens on the *MainMenu*
299 worksheet, which provides additional instruction, a report citation, and support contact
300 information (fig. 5C). The user enters the catchment characteristics summarized by StreamStats
301 into the *BasinCharacteristics* worksheet (fig. 5D) and then presses the command button to
302 compute the unregulated daily streamflows. The program then follows the process outlined in
303 figure 3 by calculating the FDC, selecting the index streamgauge using the map-correlation
304 method and transferring the timing of the streamflows at the index streamgauge to the ungauged
305 location by the QPPQ-transform method. The CRUISE-estimated streamflows are, in part,
306 computed from regional regression equations that were developed using the catchment
307 characteristics discussed in Section 3. Streamflows estimated for ungauged catchments having
308 characteristics outside the range of values used to develop the regression equations are highly
309 uncertain because these values were not used to fit the regression equations. Therefore, the

310 CRUISE tool includes a message in the *BasinCharacteristics* worksheet next to each
311 characteristic that is outside the respective ranges shown in figure 2 and detailed in Section 2.

312 The *ReferenceGaugeSelection* worksheet (fig. 5E) displays information about the
313 ungauged catchment and the selected index streamgauge, including the percent difference
314 between catchment characteristics at the ungauged and index streamgauge, the distance between
315 the between catchment characteristics at the ungauged location and index streamgauge, and the
316 estimated cross-correlation resulting from the map-correlation method. Whereas the CRUISE
317 tool automatically selects the index streamgauge estimated to be most correlated with the
318 ungauged location, the CRUISE tool also reports the five index streamgauges estimated to be
319 most correlated with the ungauged location (fig. 5E). The CRUISE tool also allows users to
320 choose from any of the potential index streamgauges in the study area if for some reason they
321 would like to select another index streamgauge, either from the five most-correlated index
322 streamgauges or another index streamgauge (fig. 5E). Users select a new index streamgauge
323 from a pull-down list and choose the update button (fig. 5E). The FDCs at the ungauged location
324 and the index streamgauge as well as the daily streamflow time series at the ungauged location
325 and at the index streamgauge are reported in cubic feet per second and cubic feet per second per
326 mile. The *ContinuousFlowDuration* worksheet (fig. 5F) displays the estimated continuous
327 exceedence probabilities, and the *ContinuousDailyFlow* worksheet (fig. 5G) displays the
328 estimated daily time series for the ungauged site. Both worksheets provide the estimated
329 streamflows in units of cubic feet per second and in cubic feet per second per square mile.

330 **5. Performance of streamflows estimated by the CRUISE model**

331 To evaluate the utility of the CRUISE tool to estimate unregulated, daily streamflow at
332 ungauged locations in the CRB, a leave-one-out cross validation for a subset of 31 study
333 streamgauges (fig. 1) was applied. These study streamgauges were selected as validation
334 streamgauges because they contained a complete period of observed record for the period
335 estimated by the CRUISE tool. Each site was removed completely from the streamflow
336 estimation process – one by one – and the parameters of the regression and variogram models
337 were re-estimated. The re-estimated parameters were then used to generate streamflow at the
338 removed site and estimated streamflows were compared with the observed streamflows.

339 Goodness of fit between observed and estimated streamflows for the entire simulation
340 period was evaluated using the Nash-Sutcliffe efficiency value [*Nash and Sutcliffe, 1970*], which
341 was computed from both the observed and estimated streamflows as well as the natural
342 logarithms of the observed and estimated streamflows (fig. 6). The natural logarithm of the
343 observed and estimated streamflows was taken to scale the daily streamflow values so that the
344 high and low streamflow values were more equally weighted in the calculation of the efficiency
345 metric. Efficiency values were mapped to determine if there was any spatial bias in the model
346 performance (fig. 7). Selected hydrographs were also plotted to visualize the interpretation of the
347 efficiency values (fig 7).

348 Percent errors in selected ecologically-relevant high- and low-flow statistics were also
349 compared (fig. 8). These streamflow statistics, as defined by *Hendrickson et al. [2006]*, include
350 measures of the magnitude of streamflows (the median monthly streamflows), duration of
351 streamflows (the 30-, 60-, and 90-day minimum and 90-day maximum streamflow), and timing
352 of streamflows (the Julian days of the 1-day minimum and maximum streamflows) (fig. 8).

353 The efficiency values in figure 6 show that the streamflows estimated by the CRUISE tool
354 generally have good agreement with the observed streamflows at the 31 validation streamgauges.
355 The minimum efficiency computed from the transformed daily streamflows is 0.69 and the
356 maximum value is 0.92 (fig. 6), with an efficiency value equal to 1 indicting perfect agreement
357 between the observed and estimated streamflows. The efficiency values for the untransformed
358 observed and estimated streamflows range from 0.04 to 0.92 (fig. 6). This decrease in efficiency
359 between the transformed and untransformed observed and estimate streamflows suggest that the
360 fit between the observed and estimated streamflows from the CRUISE tool at high streamflow
361 values is more of a challenge than the fit at the other streamflow values. Despite this, the
362 CRUISE model appears to result in high efficiency values across all validation sites (fig. 7).
363 Streamgauges in the northern portion of the CRB have lower efficiency values than
364 streamgauges in the middle and southern portions of the CRB; however, it should be noted from
365 the hydrographs in figure 7 that the CRUISE tool is able to represent the daily features of the
366 hydrographs at the validation streamgauges even though the efficiency values are relatively
367 lower in the northern portion of the study area. The selected streamflow statistics estimated by
368 the CRUISE tool also provide a reasonable match to the observed streamflow statistics at the
369 validation sites (fig. 8). The percent error for the majority of the streamflow statistics is between
370 +/- 10 percent (fig. 8), with the exception of low-flow statistics. Percent error tends to be inflated
371 for these streamflow statistics because the streamflows are already low values and, when divided
372 by the difference between the observed and estimated values, the percentages can appear high
373 even though the absolute differences between the observed and estimated streamflows values are
374 low. The efficiency values, hydrograph comparisons and flow statistics derived from the

375 observed and estimated streamflows demonstrate that the CRUISE tool can provide a reasonable
376 representation of natural streamflow time series at ungauged catchments in the CRB.

377 **6. Summary and conclusions**

378 This paper presents the Connecticut River UnImpacted Streamflow Estimation (CRUISE)
379 tool, which estimates daily, unregulated streamflow at ungauged locations in the Connecticut
380 River Basin (CRB). The CRUISE tool is freely-available and requires only an internet
381 connection and Microsoft Excel version 2003 or higher. The StreamStats web application must
382 be used to select the location of the ungauged site, delineate the catchment boundary, and
383 determine its catchment characteristics before the CRUISE tool can be used. CRUISE estimates
384 daily streamflow time series for a 44-year period of record from October 1, 1960 through
385 September 30, 2004. Daily streamflow is estimated by a three-part process: 1) estimation of the
386 daily, period-of-record flow-duration curve at the ungauged location, 2) selection of an index
387 streamgauge, and 3) use of the index streamgauge to transfer the flow-duration curve to a time
388 series of daily streamflow. The CRUISE tool provided reliable estimates of observed daily
389 streamflows at 31 validation streamgauges across the CRB. The coupling of the StreamStats and
390 CRUISE tools presents a modeling and software framework that can be used to develop point-
391 and-click, GIS-based, daily-streamflow estimates needed for water management decisions at
392 ungauged stream locations for other regions.

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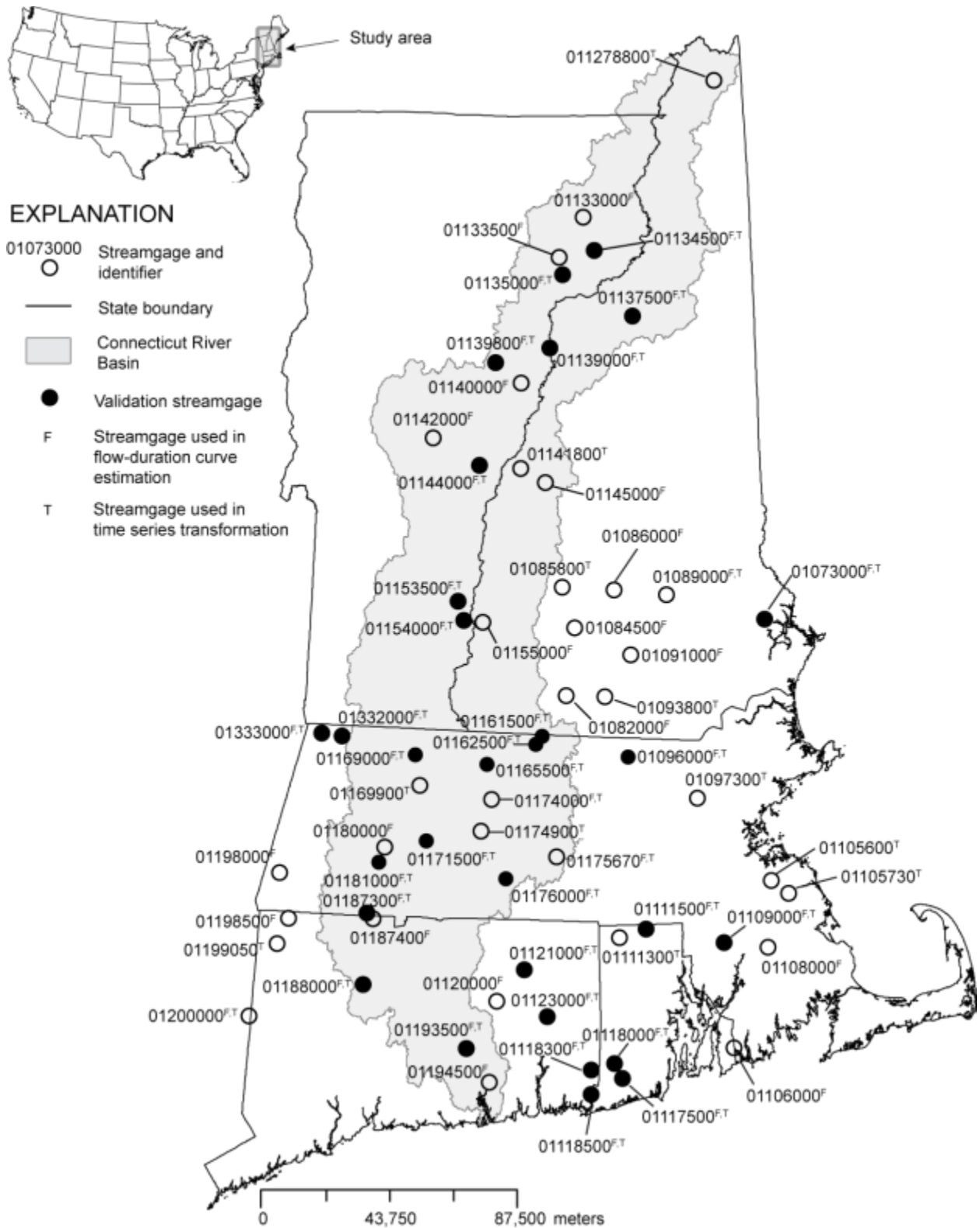
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401 the data for the CRUISE tool, and Scott Olsen of the U.S. Geological Survey, who provided a list
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405 **References**

- 406 Archfield, S. A., and R. M. Vogel, 2010. Map correlation method: Selection of a reference
407 streamgauge to estimate daily streamflow at ungaged catchments, *Water Resour. Res.*, 46,
408 W10513, doi:10.1029/2009WR008481.
- 409 Archfield, S., R. Vogel, P. Steeves, S. Brandt, P. Weiskel, and S. Garabedian, 2010. The
410 Massachusetts Sustainable-Yield Estimator: A decision-support tool to assess water
411 availability at ungaged sites in Massachusetts, U.S. Geological Survey Scientific
412 Investigations Report 2009-5227, 41 p. plus CD-ROM.
- 413 Armstrong, D. S., G. W. Parker, and T. A. Richards, 2008. Characteristics and classification of
414 least altered streamflows in Massachusetts, U.S. Geological Survey Scientific
415 Investigations Report, 20075291, 113 p. plus CD-ROM.
- 416 Duan, N., 1983) Smearing estimate—a nonparametric retransformation method, *J. Am. Stat.*
417 *Assoc.*, 78, 383, pp. 605-610.
- 418 Castellarin, A., G. Galeati, L. Brandimarte, A. Montanari, and A. Brath, 2004. Regional flow-
419 duration curves: reliability for ungaged basins, *Adv. Water Resour.*, 27, 10, 953-965
- 420 ESRI, Inc., 2009. Arc-Hydro Tools - Tutorial, Version 1.3 - January 2009, ESRI, Inc., Redlands,
421 CA, available at http://andersonruhoff.googlepages.com/ArcHydro_Tutorial.pdf.
- 422 Falcone, J. A., D. M. Carlisle, D. M. Wolock, and M. R. Meador, 2010. GAGES: A stream gage
423 database for evaluating natural and altered flow conditions in the coterminous United
424 States, *Ecology*, 91, 612.
- 425 Fennessey, N. M., 1994. A hydro-climatological model of daily streamflow for the northeast
426 United States, Ph.D. dissertation, Tufts University, Department of Civil and
427 Environmental Engineering.

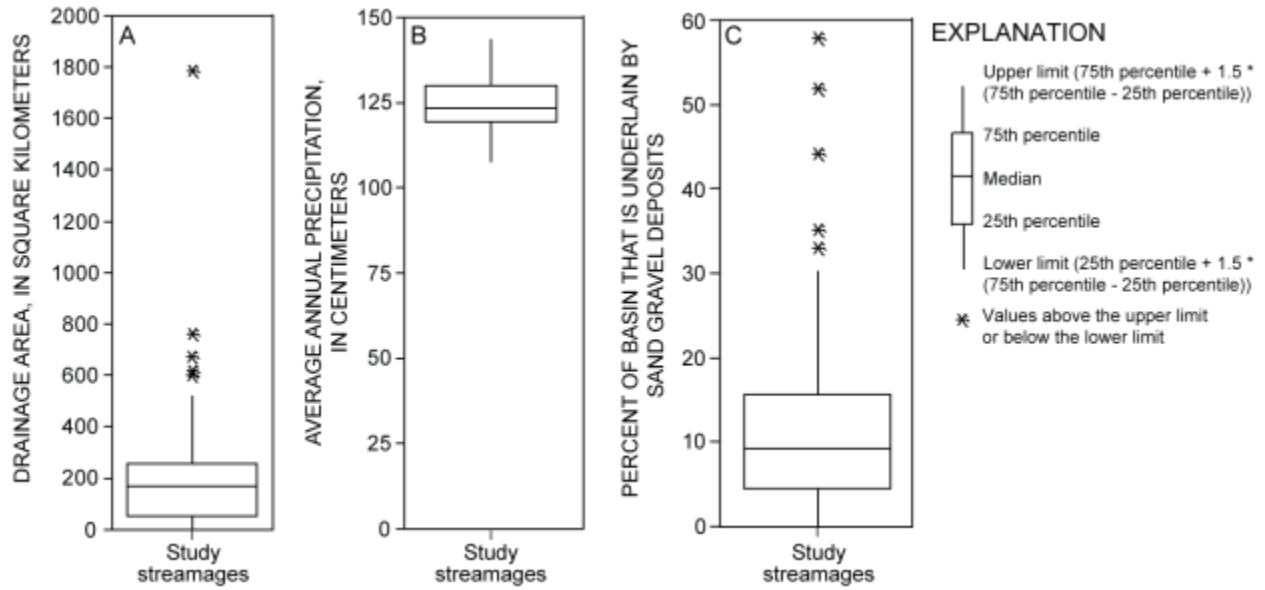
- 428 Granato, G.E., 2009. Computer programs for obtaining and analyzing daily mean streamflow
429 data from the U.S. Geological Survey National Water Information System Web Site, U.S.
430 Geological Survey Open-File Report 2008–1362, 123 p.
- 431 Helsel, D., and R. Hirsch, 2002. Statistical Methods in Water Resources Techniques of Water
432 Resources Investigations, Book 4, Chapter A3, U.S. Geological Survey.
- 433 Henriksen, J.A., J. Heasley, J.G. Kennen, and S. Nieswand, 2006. User’s manual for the
434 Hydroecological Integrity Assessment Process software, U.S. Geological Survey Open-
435 File Report 2006–1093, 71 p.
- 436 Hirsch, R., 1979. Evaluation of some record reconstruction techniques, *Water Resour. Res.*, 15,
437 6, 1781-1790, ISSN 0043-1397.
- 438 Holtschlag, D.J., 2009. Application guide for AFINCH, analysis of flows in networks of
439 channels) described by NHDPlus, U.S. Geological Survey Scientific Investigations
440 Report 2009-5188, 106 p.
- 441 Hughes, D.A., and V.U. Smakhtin, 1996. Daily flow time series patching or extension: a spatial
442 interpolation approach based on flow duration curves: *Hydrolog. Sci. J.*, 41, 6, 851–871.
- 443 Isaaks, E. H., and R. M. Srivastava, 1989. *An Introduction to Applied Geostatistics*, first ed.,
444 Oxford University Press, New York.
- 445 Mahoamoud, Y. M., 2008. Prediction of daily flow duration curves and streamflow for ungauged
446 catchments using regional flow duration curves, *Hydrolog. Sci. J.*, 53, 4, 706-724.
- 447 Nash, J. E., and J. V. Sutcliffe, 1970. River flow forecasting through conceptual models part I - a
448 discussion of principles, *J. of Hydrol.*, 10, 3, 282–290.
- 449 Poff, N.L., J.D. Allen, M.B. Bain, J.R. Karr, K.L. Prestagard, B.D. Richter, R.E. Sparks, and
450 J.C. Stromberg, 1997. The natural-flow regime—A paradigm for river conservation and
451 restoration: *Bioscience*, 47, 769–784.
- 452 Ribeiro Jr., P., and P. Diggle, 2001. *gEO*: A package for geostatistical analysis, *R-News*, 1, 2.
- 453 Ries, K. G., III; Guthrie, J. G.; Rea, A. H.; Steeves, P. A.; Stewart, D. W., 2008. StreamStats: A
454 Water Resources Web Application, U.S. Geological Survey Fact Sheet 2008-3067, 6 p.,
455 available on line at <http://pubs.usgs.gov/fs/2008/3067/>.
- 456 Ries, K.G., III, and P.J. Friesz, 2000. Methods for estimating low-flow statistics for
457 Massachusetts Streams, U.S. Geological Survey Water-Resources Investigations Report
458 2000-4135, 81 p., available on line at <http://pubs.usgs.gov/wri/wri004135/>.
- 459 Skøien, J. O., and G. Blöschl, 2007. Spatiotemporal topological kriging of runoff time series,
460 *Water Resour. Res.*, 43, 9, doi:10.1029/2006WR005760.
- 461 Smakhtin, V. U., 1999. Generation of natural daily flow time-series in regulated rivers using a
462 non-linear spatial interpolation technique, *Regul. Rivers: Res. Mgmt*, 15, 311-323.

- 463 Smakhtin V.U. and N. Eriyagama, 2008. Developing a software package for global desktop
464 assessment of environmental flows. *Environ. Model. Softw.* 23, 12, December 2008.
465 1396-1406. doi:10.1016/j.envsoft.2008.04.002.
- 466 Stedinger, J.R., R.M. Vogel, and E. Foufoula-Georgiou, 1993. Frequency analysis of extreme
467 events, Chapter 18, in Maidment, D.R., ed., *Handbook of Hydrology*: McGraw-Hill Book
468 Company, New York.
- 469 Vogel, R.M., and N.M. Fennessey, 1994. Flow duration curves I: A new interpretation and
470 confidence intervals, *J. Water Res. Pl.-ASCE*, 120,4, p. 485-504.
- 471 Vogel, R.M., and J.R. Stedinger, 1985. Minimum variance streamflow record augmentation
472 procedures, *Water Resour. Res.*, 21,5, p. 715–723.
- 473 Williamson, T.N., K.R. Odom, J.K. Newson, A.C. Downs, H.L. Nelson Jr., P.J. Cinotto and
474 M.A. Ayers, 2009. The Water Availability Tool for Environmental Resources,
475 WATER)—A water-budget modeling approach for managing water-supply resources in
476 Kentucky—Phase I—Data processing, model development, and application to non-karst
477 areas: U.S. Geological Survey Scientific Investigations Report 2009–5248, 34 p.
- 478 Zhang, Y., and F. H. S. Chiew, 2009. Relative merits of different methods for runoff predictions
479 in ungauged catchments, *Water Resour. Res.*, 45, W07412, doi:10.1029/2008WR007504.
- 480



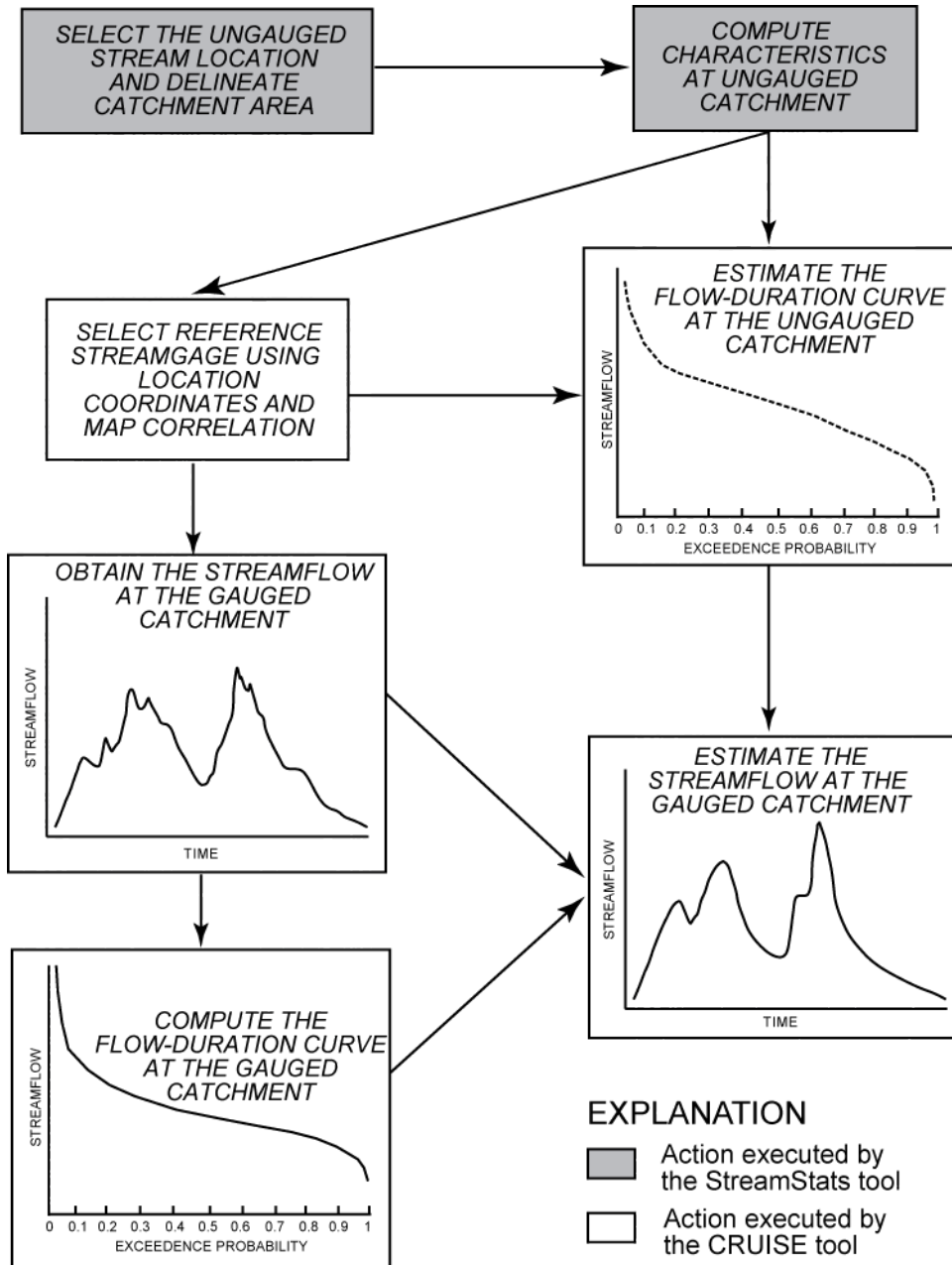
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482 Figure 1. Map showing the locations of the streamgauges used to estimate unregulated, daily
 483 streamflow at ungauged locations in the Connecticut River Basin, northeast United States.



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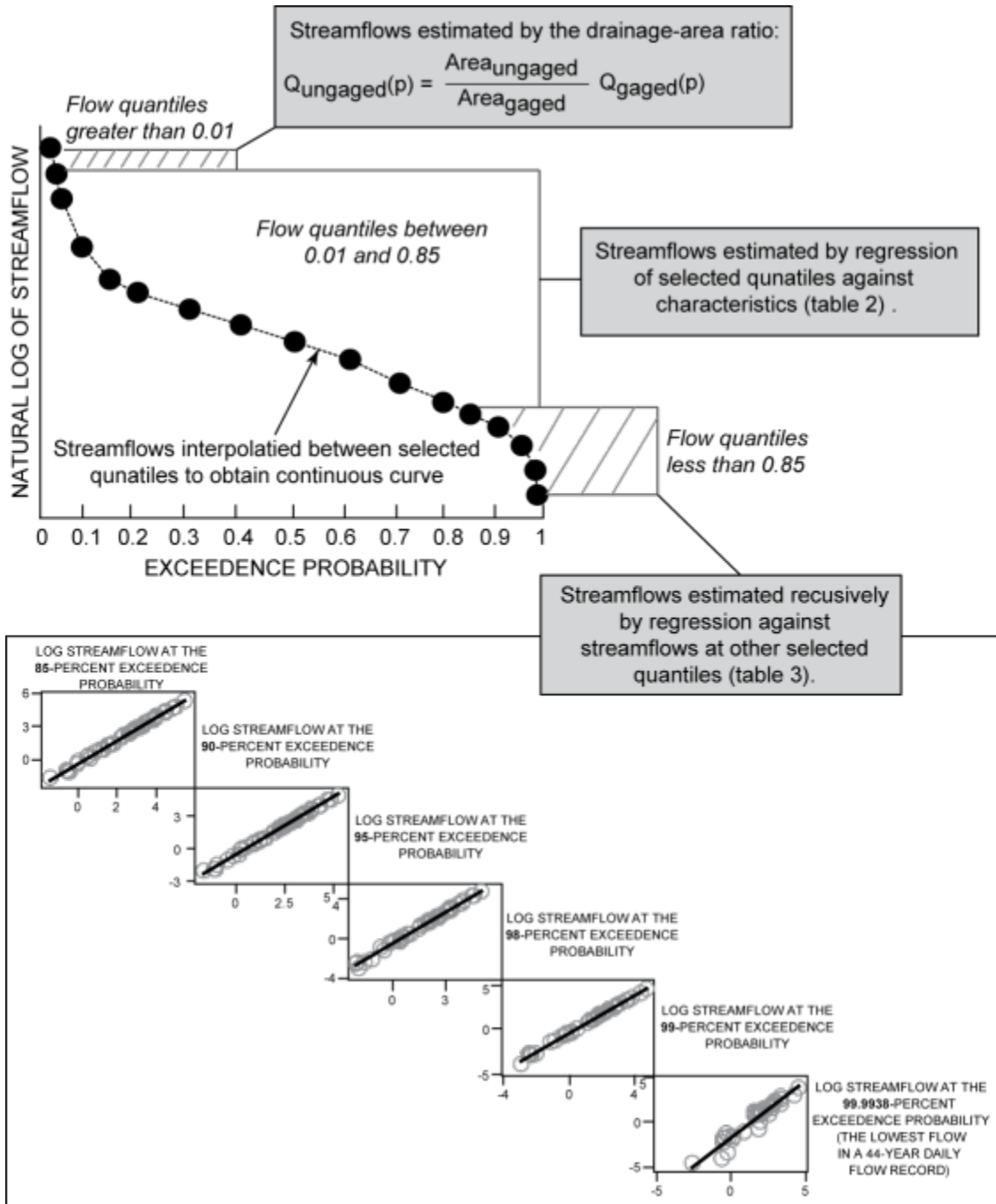
Figure 2. Boxplots of the catchment characteristics used to estimate unregulated, daily streamflow at ungauged locations in the Connecticut River Basin, northeast United States and the subset of catchment characteristics used to validate the streamflow estimates.



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Figure 3. Diagram of the process to estimate unregulated, daily streamflow at ungauged locations.

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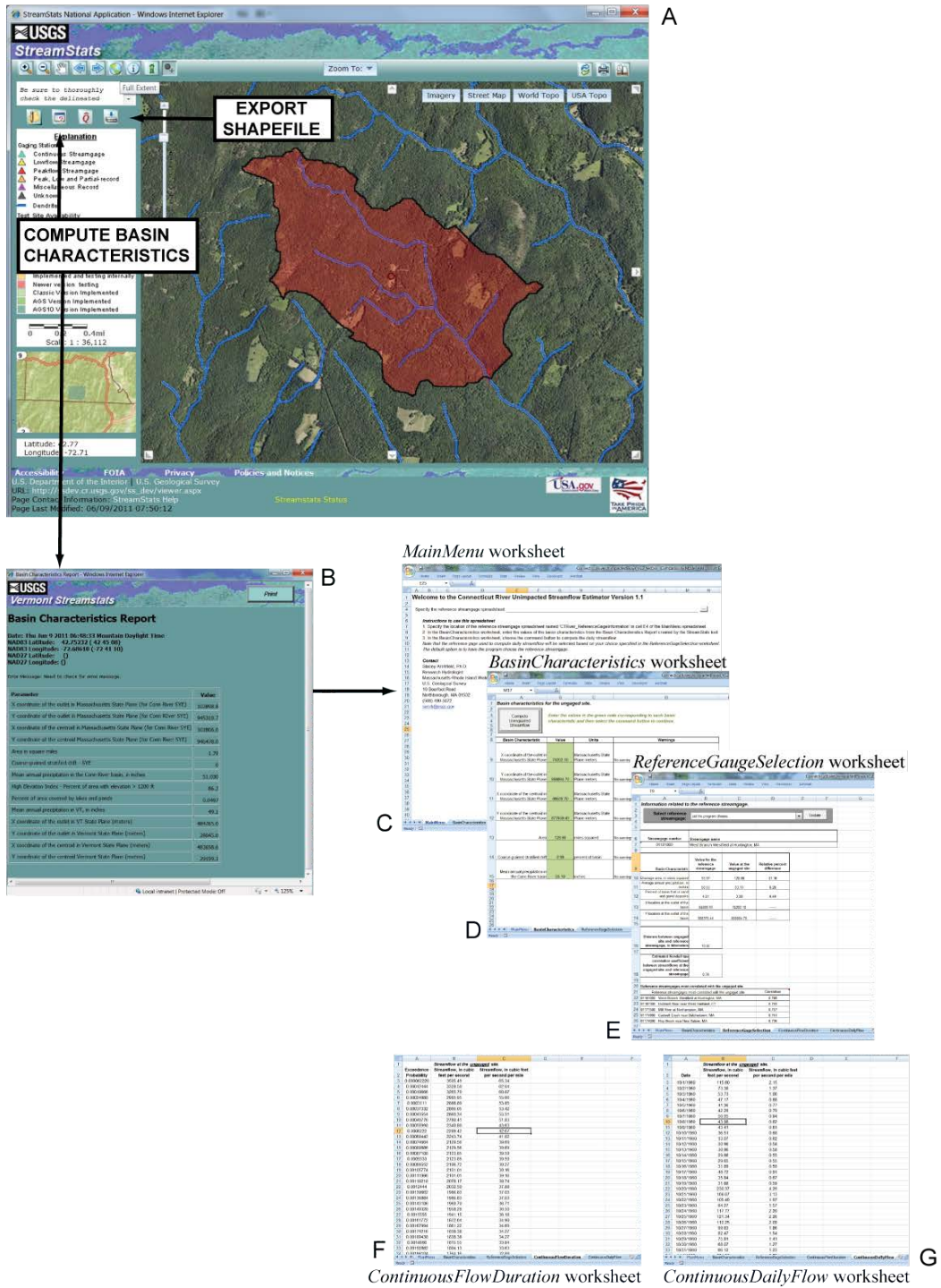


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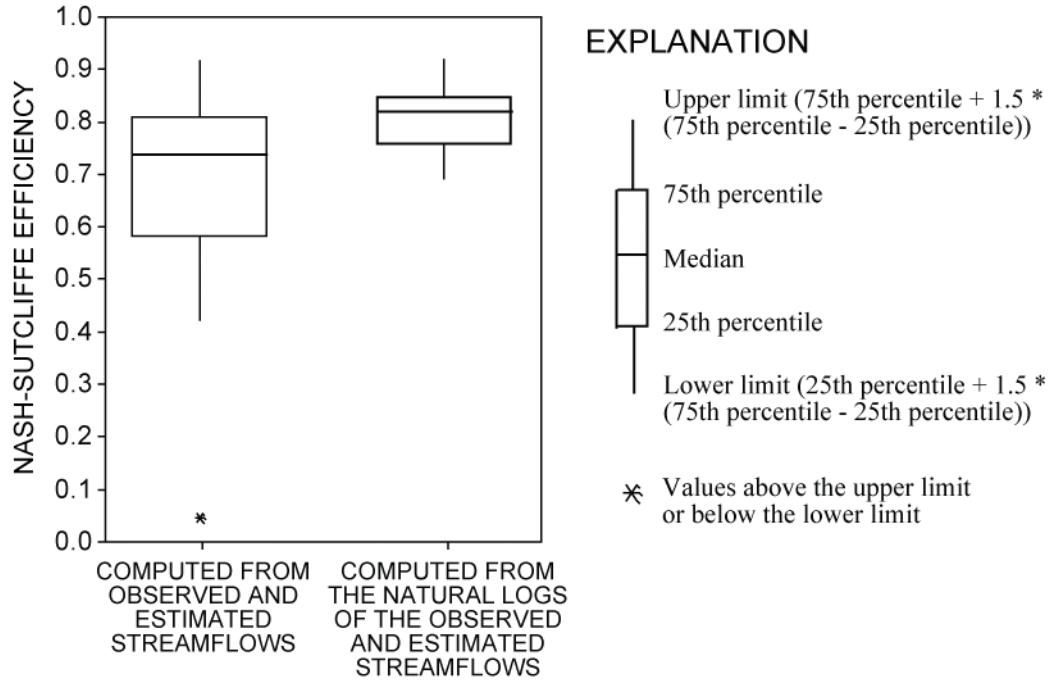
496 Figure 4. Diagram showing the methods used to estimate a continuous, daily flow duration at an
 497 ungaged location.

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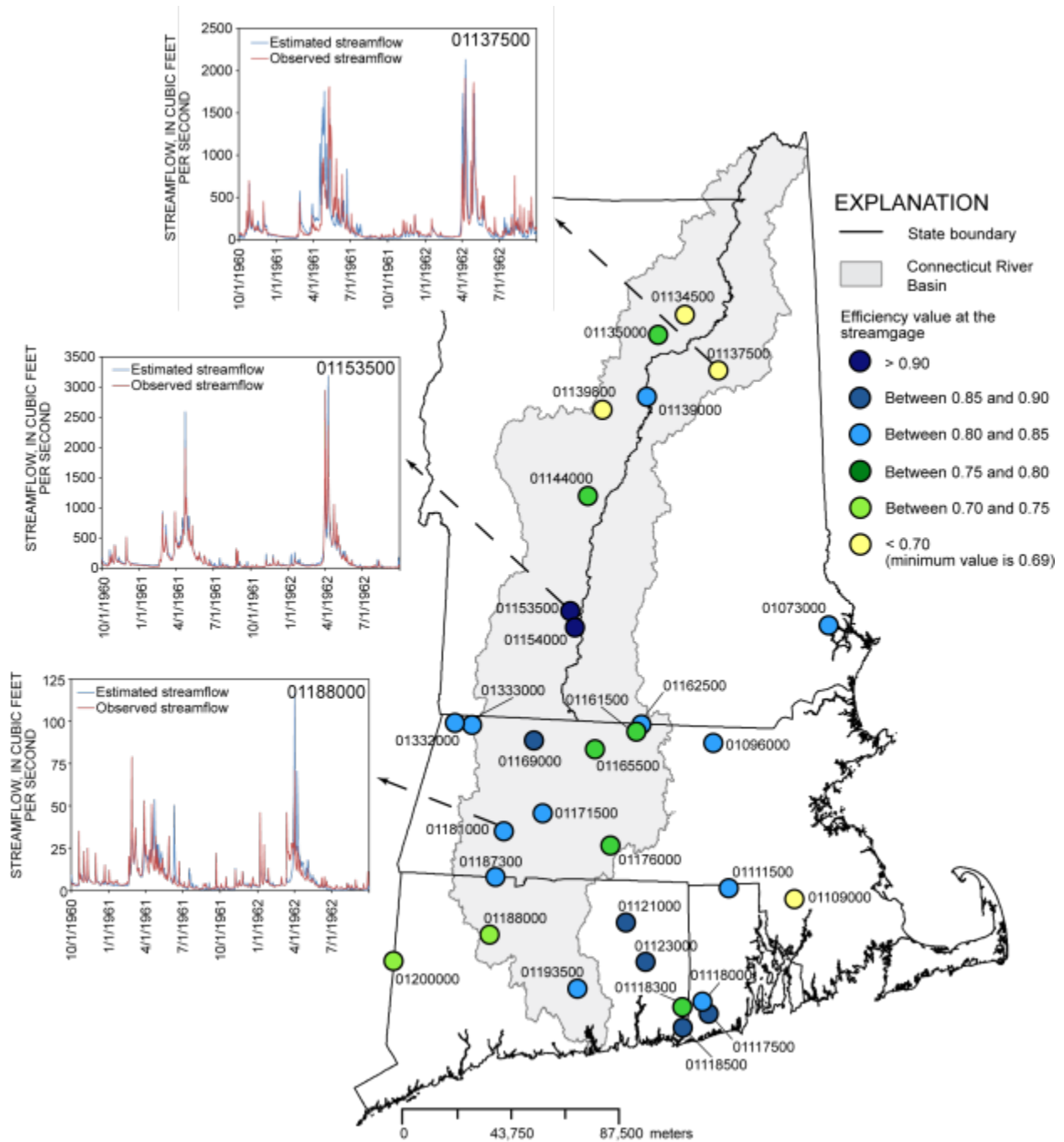
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Figure 5. Screen captures showing the decision-support tool used to estimate daily, unregulated time series in the Connecticut River Basin. The program delineates a catchment for the ungauged location selected by the user (A) and summarizes the catchment characteristics (B). The user then inputs these characteristics into a spreadsheet program (C-E) that generates the daily, period of record flow-duration curve (F) and the daily streamflow time series (G).

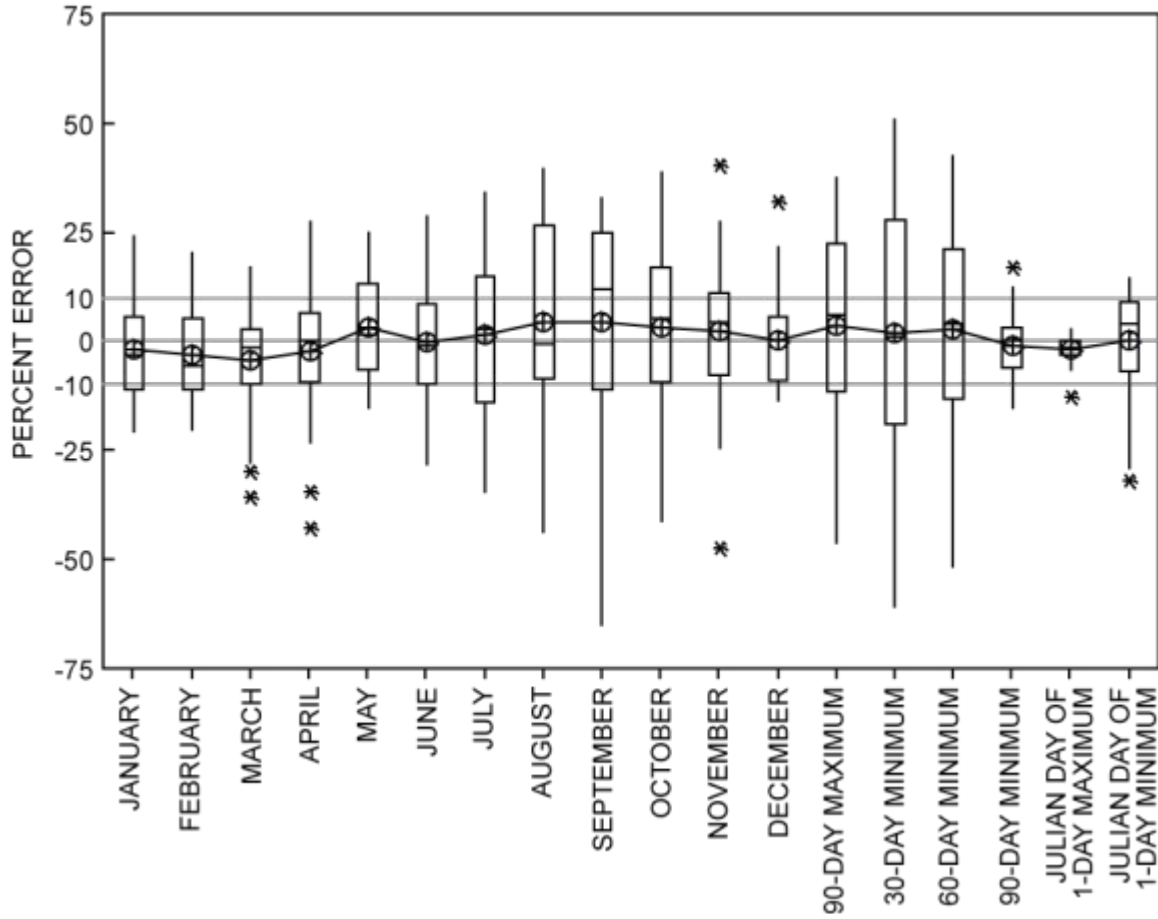


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Figure 6. Range of efficiency values computed between the observed and estimated streamflows at the 31 validation streamgauges.



510
 511 Figure 7. Spatial distribution of efficiency values resulting from log-transformed observed and
 512 estimated daily streamflow at 31 validation streamgages and selected hydrographs of observed
 513 and estimated streamflow for the period from October 1, 1960 through September 30, 1962.
 514



515
 516 | Figure 8. Percent error between observed and estimated monthly mean flows and other
 517 | ecologically-relevant streamflow statistics at the 31 validation streamgauges.
 518

519

520 Table 1. List of streamgauges used to estimate unregulated, daily streamflow at ungauged

521 locations in the Connecticut River Basin.

Station Number	Station name	Period of record
01073000	Oyster River near Durham, NH	December 15, 1934 - December 31, 2004
01082000	Contocook River at Peterborough, NH	July 7, 1945 - September 30, 1977
01084500	Beard Brook near Hillsboro, NH	October 1, 1945 - September 30, 1970
01085800	West Branch Warner River near Bradford, NH	May 22, 1962 - September 30, 2004
01086000	Warner River at Davisville, NH	October 1, 1939 - September 30, 1978
01089000	Soucook River near Concord, NH	October 1, 1951 - September 30, 1987
01091000	South Branch Piscataquog River near Goffstown, NH	July 27, 1940 - September 30, 1978
01093800	Stony Brook tributary near Temple, NH	May 1, 1963 - September 30, 2004
01096000	Squannacook River near West Groton, MA	October 1, 1949 - December 31, 2004
01097300	Nashoba Brook near Acton, MA	July 26, 1963 - December 31, 2004
01105600	Old Swamp River near South Weymouth, MA	May 20, 1966 - July 24, 2006
01105730	Indian Head River at Hanover, MA	July 8, 1966 - July 24, 2006
01106000	Adamsville Brook at Adamsville, RI	October 1, 1940 - September 30, 1978
01108000	Taunton River near Bridgewater, MA	October 1, 1929 - April 23, 1976
01109000	Wading River near Norton, MA	June 1, 1925 - December 31, 2004
01111300	Nipmuc River near Harrisville, RI	March 1, 1964 - September 30, 1991
01111500	Branch Riverb at Forestdale, RI	January 24, 1940 - December 31, 2004
01117500	Pawcatuck River at Wood River Junction, RI	December 7, 1940 - December 31, 2004
01118000	Wood River Hope Valley, RI	March 12, 1941 - December 31, 2004
01118300	Pendleton Hill Brook near Clarks Falls, CT	October 1, 1958 - December 31, 2004
01118500	Pawtucket River at Westerly, RI	November 27, 1940 - December 31, 2004
01120000	Hop Brook near Columbia, CT	October 1, 1932 - October 6, 1971
01121000	Mount Hope River near Warrenville, CT	October 1, 1940 - December 31, 2004
01123000	Little River near Hanover, CT	October 1, 1951 - December 31, 2004
01127880	Big Brook Near Pittsburg Nh	December 1, 1963 - January 1, 1984
01133000	East Branch Passumpsic River near East Haven, VT	October 1, 1948 - September 1, 1979
01133500	Passumpsic River near St. Johnsbury, VT	May 1, 1909 - July 1, 1919
01134500	Moose River at Victory, VT	January 1, 1947 - May 12, 2010
01135000	Moose River at St. Johnsbury, VT	August 1, 1928 - September 1, 1983
01137500	Ammonoosuc River at Bethlehem Junction, NH	August 1, 1939 - May 12, 2010
01139000	Wells River at Wells River, VT	August 1, 1940 - May 12, 2010
01139800	East Orange Branch at East Orange, VT	June 1, 1958 - May 12, 2010
01140000	South Branch Waits River near Bradford, VT	April 1, 1940 - September 1, 1951
01141800	Mink Brook near Etna, NH	August 1, 1962 - September 1, 1998
01142000	White River near Bethel, VT	June 1, 1931 - September 1, 1955
01144000	White River at West Hartford, VT	October 1, 1951 - May 12, 2010
01145000	Mascoma River at West Canaan, NH	July 1, 1939 - September 1, 1978
01153500	Williams River near Rockingham, VT	June 1, 1940 - September 1, 1984
01154000	Saxtons River at Saxtons River, VT	June 20, 1940 - September 30, 1982
01155000	Cold River at Drewsville, NH	June 23, 1940 - September 30, 1978
01161500	Tarbell Brook near Winchendon, MA	May 29, 1916 - September 6, 1983
01162500	Priest Brook near Winchendon, MA	October 1, 1936 - December 31, 2004
01165500	Moss Brook at Wendell Depot, MA	June 1, 1916 - September 30, 1982
01169000	North River at Shattuckville, MA	December 13, 1939 - December 31, 2004
01169900	South River near Conway, MA	January 1, 1967 - December 31, 2004
01171500	Mill River at Northampton, MA	November 18, 1938 - December 31, 2004
01174000	Hop Brook near New Salem, MA	November 19, 1947 - September 30, 1982
01174900	Cadwell Creek near Belchertown, MA	July 13, 1961 - September 30, 1997
01175670	Sevenmile River near Spencer, MA	December 1, 1960 - December 31, 2004
01176000	Quaboag River at West Brimfield, MA	August 19, 1912 - December 31, 2004
01180000	Sykes Brook at Knightville, MA	June 20, 1945 - July 18, 1974
01181000	West Branch Westfield at Huntington, MA	September 1, 1935 - December 31, 2004

01187300	Hubbard River near West Hartland, CT	August 4, 1959 - December 31, 2004
01187400	Valley Brook near West Hartland, CT	October 1, 1940 - September 30, 1972
01188000	Burlington Brook near Burlington, CT	October 1, 1931 - December 31, 2004
01193500	Salmon River near East Hampton, CT	October 1, 1928 - December 31, 2004
01194500	East Branch Eightmile River near North Lyme, CT	October 1, 1937 - October 6, 1981
01198000	Green River near Great Barrington, MA	October 1, 1951 - September 30, 1971
01198500	Blackberry River at Canaan, CT	October 1, 1949 - October 20, 1971
01199050	Salmon Creek at Lime Rock, CT	October 1, 1961 - December 31, 2004
01200000	Ten Mile River, CT	October 1, 1930 - April 4, 1988
01332000	North Branch Hoosic River at North Adams, MA	June 22, 1931 - September 30, 1990
01333000	Green River at Williamstown, MA	September 20, 1949 - December 31, 2004

522

523 Table 2. Number of streamgages, goodness of fit values, explanatory variables, and estimated
 524 regression parameters for streamflows estimated from catchment characteristics.

525
 526 [%RMSE, Percent root-mean square error; **, parameters not included in regression equation; †,
 527 Bias correction factor computed from Duan (1983)]
 528

Exceedence probability	General regression information			Characteristics in the regression equation and coefficient value						
	Number of streamgages used to develop regression equation	%RMSE	Efficiency value	Constant term	Drainage area	Average annual precipitation	Percent of basin that is underlain by sand and gravel deposits	Y-location of the basin centroid	X-location of the basin centroid	Bias correlation factor†
0.02	51	1.49	0.99	-26.57576	0.95898668	2.32615049	**	1.44624683	**	1.01034136
0.05	51	0.62	1.00	-19.31477	0.97753613	1.75206249	**	1.04571137	**	1.00231622
0.1	51	0.73	0.99	-2.122368	0.99821922	0.91063827	**	**	**	1.00149376
0.15	51	0.60	1.00	-2.977718	1.0050255	1.05886106	**	**	**	0.99718242
0.2	51	0.86	0.99	-3.693535	1.00370636	1.1919636	**	**	**	0.99567544
0.25	51	1.32	0.98	-4.668431	1.01096501	1.38901861	**	**	**	0.99499049
0.3	51	1.86	0.98	-5.539372	1.01366908	1.56884624	**	**	**	0.99497319
0.4	51	3.00	0.96	-6.759127	1.0206135	1.79999061	**	**	**	0.99601846
0.5	51	3.86	0.95	-7.680269	1.02689366	1.95768173	**	**	**	0.99821139
0.6	50	4.40	0.96	-8.346613	1.0184328	2.01229612	0.080378539	**	**	1.01839085
0.7	50	6.61	0.94	-8.449954	1.04799744	1.90718911	0.094903615	**	**	1.02784995
0.75	50	9.24	0.93	-8.745009	1.06545026	1.90731506	0.10398442	**	**	1.02430492
0.8	50	13.58	0.92	-9.108501	1.09514097	1.9007584	0.125122097	**	**	1.0379095
0.85	50	21.20	0.90	-9.315441	1.12388596	1.8479745	0.151546518	**	**	1.05647159

529

530 Table 3. Number of streamgages, goodness of fit values, explanatory variables, and estimated
 531 regression parameters for streamflows estimated from other streamflow quantiles.

532
 533 [%RMSE, Percent root-mean square error; †, Bias correction factor computed from Duan
 534 (1983)]
 535

Exceedence probability	General regression information			Characteristics in the regression equation and coefficient value			
	Number of streamgages used to develop regression equation	%RMSE	Efficiency value	Constant term	Coefficient on explanatory variable	Explanatory variable	Bias correlation factor†
0.9	50	32.36	0.89	-0.4112	1.0511	Streamflow at the 0.85 exceedence probability	1.0004
0.95	50	57.15	0.85	-0.4991	1.0607	Streamflow at the 0.9 exceedence probability	0.9986
0.98	50	67.36	0.79	-0.4695	1.0567	Streamflow at the 0.95 exceedence probability	1.0103
0.99	50	102.33	0.71	-0.3011	1.0467	Streamflow at the 0.98 exceedence probability	1.0000
0.999938	34	825.08	-1.30	-1.6658	1.2826	Streamflow at the 0.99 exceedence probability	1.2011

536

537 Table 4. Variogram model parameters and root-mean-square error value resulting from a leave-
 538 one-out cross validation of the variogram models.
 539

Station Number	Variance parameter	Range parameter	Root-mean-square error
01073000	0.0411	697945.4362	0.0399
01085800	0.0115	267272.8077	0.0388
01089000	0.0112	269793.6063	0.0462
01093800	0.0147	267272.7273	0.0416
01096000	0.0389	607472.9297	0.0469
01097300	0.0261	374218.0554	0.0488
01105600	0.0621	557922.7912	0.0488
01105730	0.0677	547625.3299	0.0447
01109000	0.0588	489036.3840	0.0487
01111300	0.0444	435141.4397	0.0470
01111500	0.0649	664951.4696	0.0452
01117500	0.0964	846131.5260	0.0548
01118000	0.0680	547336.8809	0.0456
01118300	0.0541	478962.6030	0.0421
01118500	0.1548	1255724.6703	0.0469
01121000	0.0440	467562.3777	0.0442
01123000	0.0487	476803.1943	0.0457
01127880	0.0475	451474.0307	0.0241
01134500	0.0585	593052.1148	0.0491
01135000	0.0828	885228.5293	0.0574
01137500	0.0421	469510.7730	0.0194
01139000	0.0354	483627.8140	0.0309
01139800	0.0224	369057.2000	0.0255
01141800	0.0116	267272.7273	0.0264
01144000	0.0155	302281.0433	0.0328
01153500	0.0135	267272.7081	0.0409
01154000	0.0129	213818.1818	0.0470
01161500	0.0187	337256.6753	0.0447
01162500	0.0176	291135.1932	0.0436
01165500	0.0291	445510.0450	0.0417
01169000	0.0190	317944.4643	0.0402
01169900	0.0245	398758.9250	0.0442
01171500	0.0310	393869.0688	0.0454
01174000	0.0249	330495.4703	0.0443
01174900	0.0321	412573.1453	0.0430
01175670	0.0366	486730.2368	0.0463
01176000	0.0357	526274.7021	0.0498
01181000	0.0333	502453.4839	0.0426
01187300	0.0566	846080.6046	0.0422
01188000	0.0313	454196.0564	0.0427
01193500	0.0412	435477.5668	0.0445
01199050	0.0212	368184.1116	0.0414
01200000	0.0401	538909.4325	0.0444
01332000	0.0114	175180.2029	0.0370
01333000	0.0148	267272.7273	0.0341

540